

The hydroclimate of Montana over the past half-century: comparison against global trends

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bit.ly/StoyGScholar



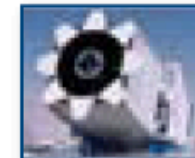
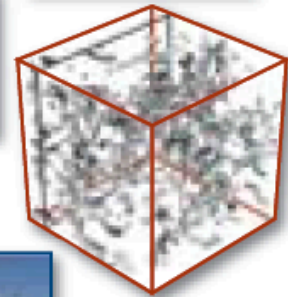
Motivation:

Science Paradigms

- Thousand years ago:
science was **empirical**
describing natural phenomena
- Last few hundred years:
theoretical branch
using models, generalizations
- Last few decades:
a **computational** branch
simulating complex phenomena
- Today: **data exploration** (eScience)
unify theory, experiment, and simulation
 - Data captured by instruments
or generated by simulator
 - Processed by software
 - Information/knowledge stored in computer
 - Scientist analyzes database/files
using data management and statistics



$$\left(\frac{\dot{a}}{a}\right)^2 = \frac{4\pi G\rho}{3} - K\frac{c^2}{a^2}$$

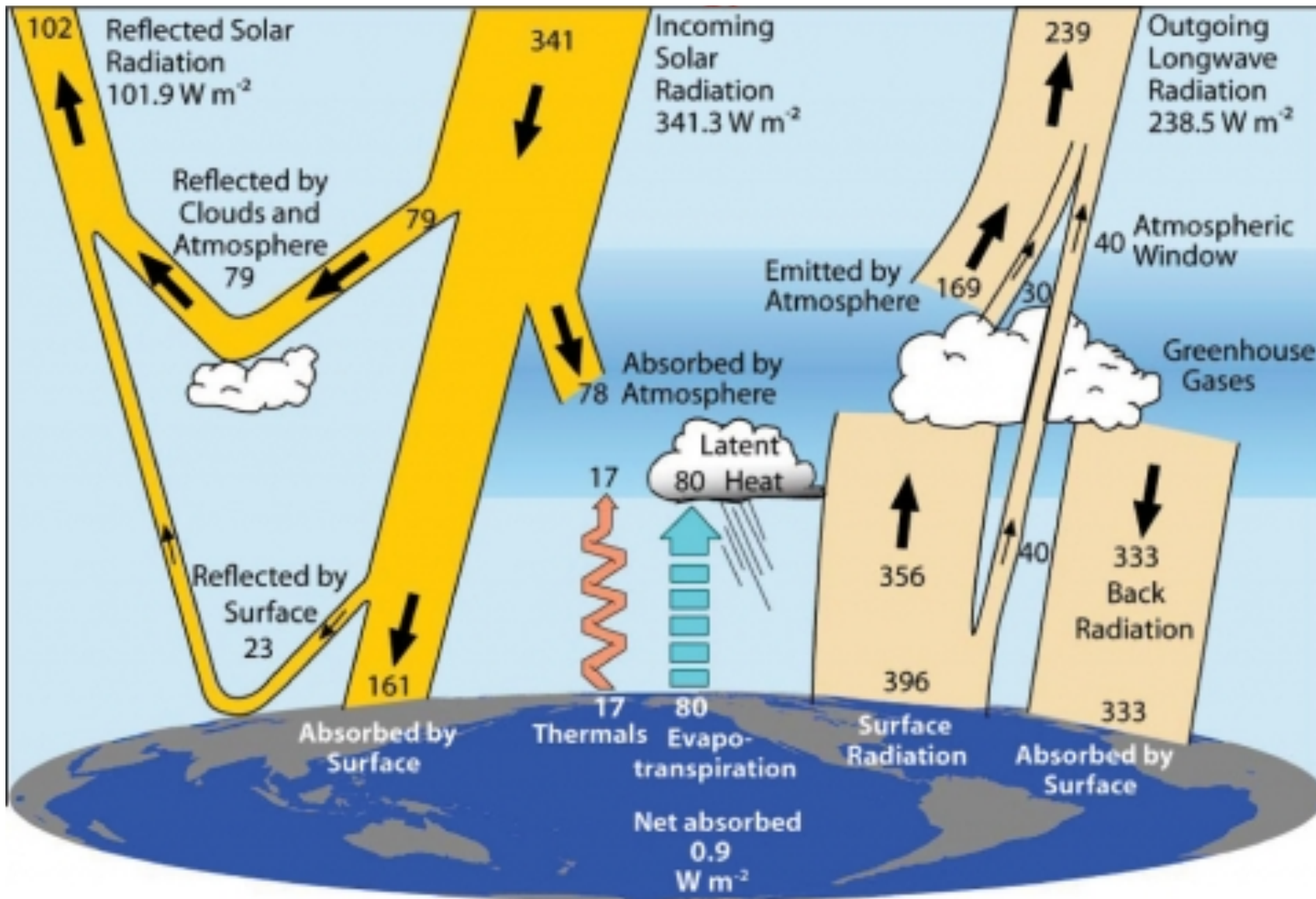


- Gray (2007)

IPCC WG1 AR5 just released...but what about MT?

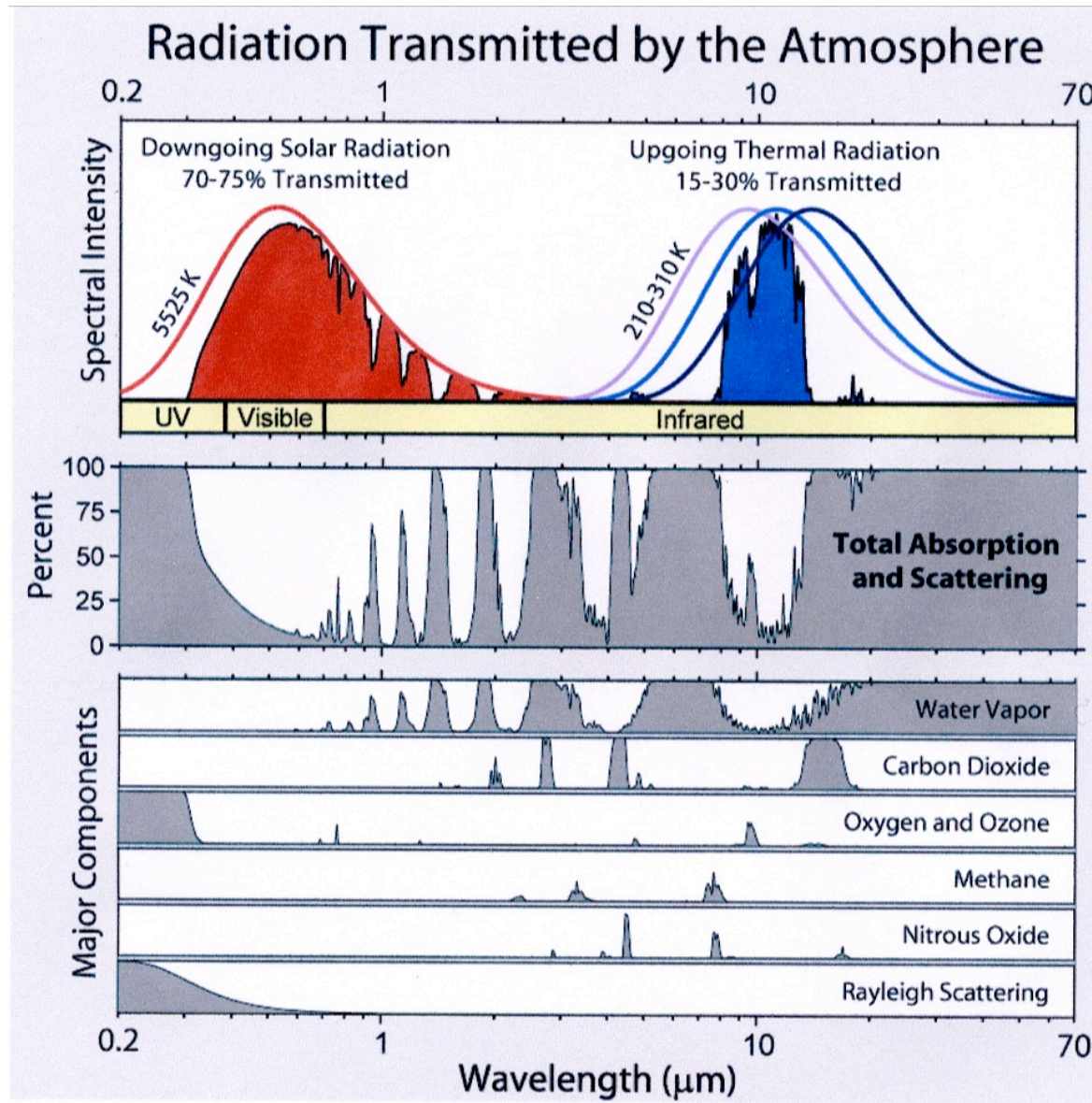


Global Energy Balance (W m^{-2})

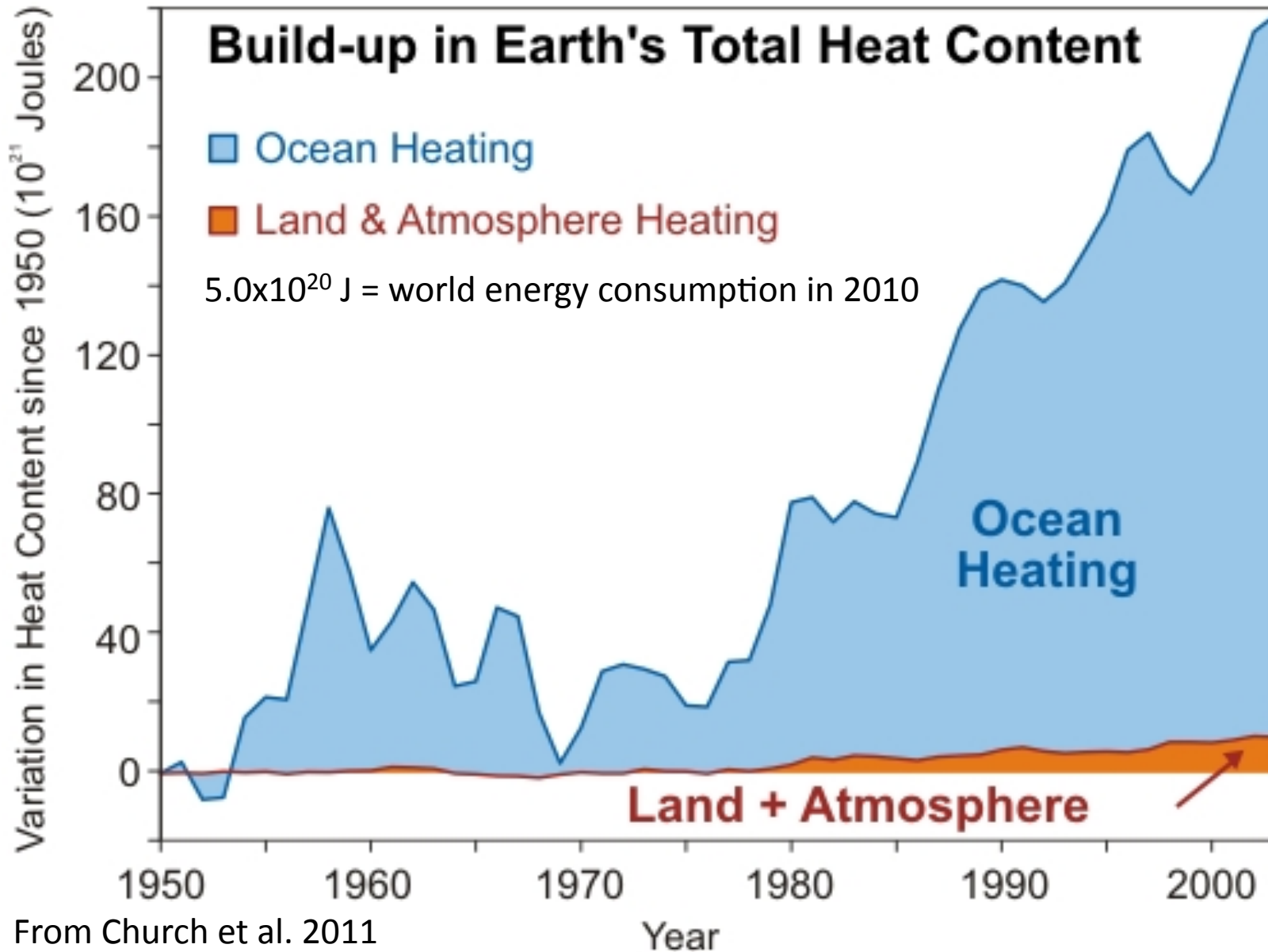


-Ultimately following Kiehl and Trenberth (1997) *BAMS* 78:197-208.

Changes to atmospheric transmissivity in the infrared



Global Change vs. Global Warming



What does it all mean for water resources in MT?



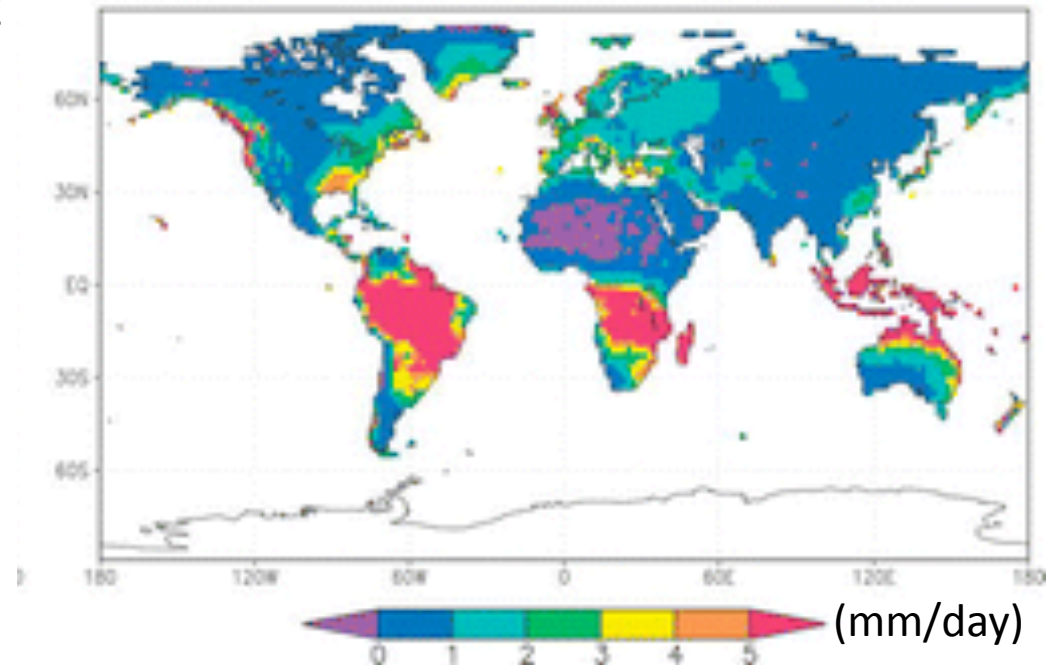
Exploring the Sheffield et al. (2006) meteorological forcing dataset

Daily global 0.5° data product

Designed for hydrology:

- Precipitation
- Air temperature
- Specific Humidity
- Air Pressure
- Wind speed
- Incident shortwave radiation
- Downwelling longwave radiation

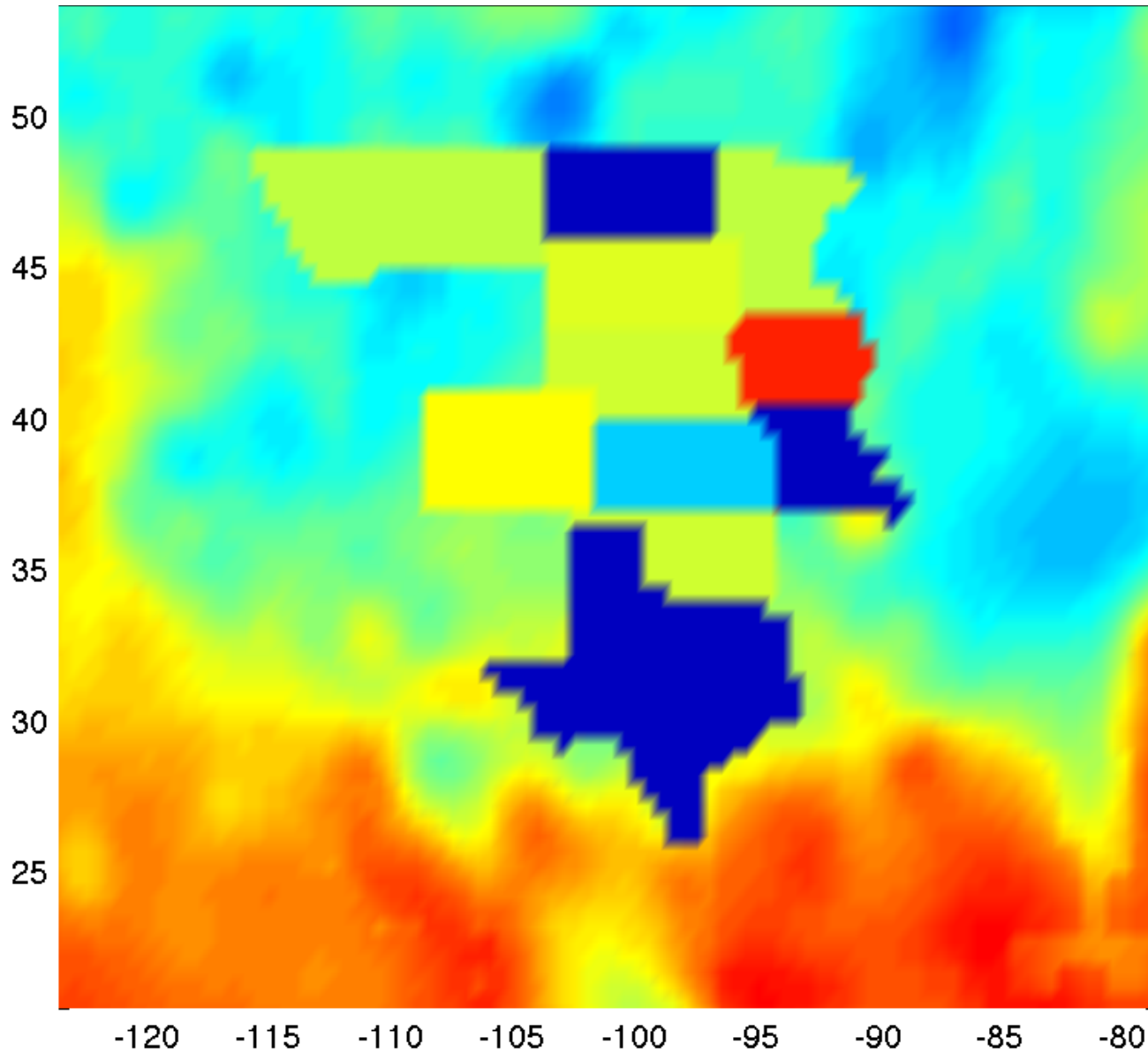
(d) NCEP Corrected January Precip.



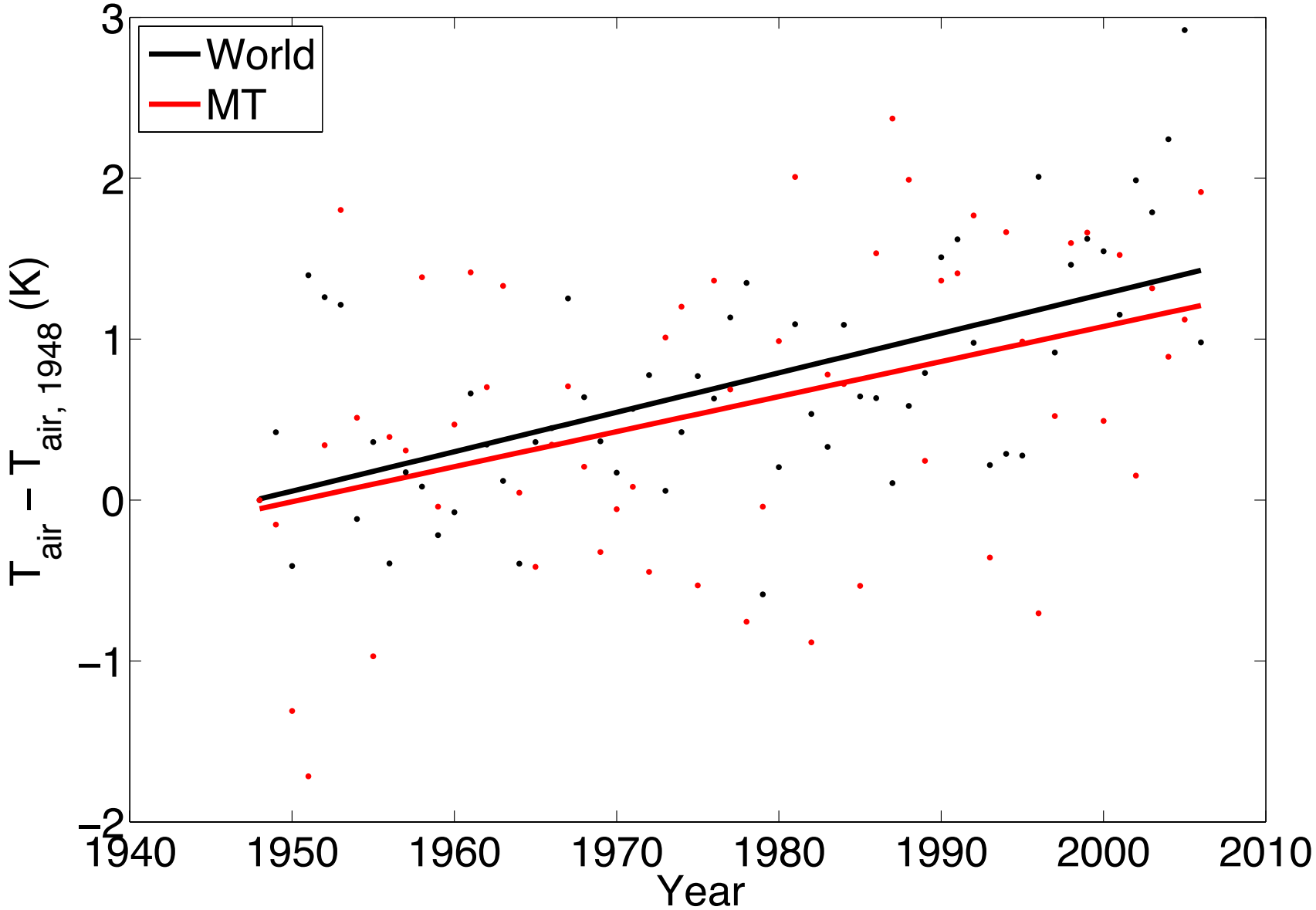
Combination of observational and reanalysis data products

Bias correction for precipitation gage undercatch

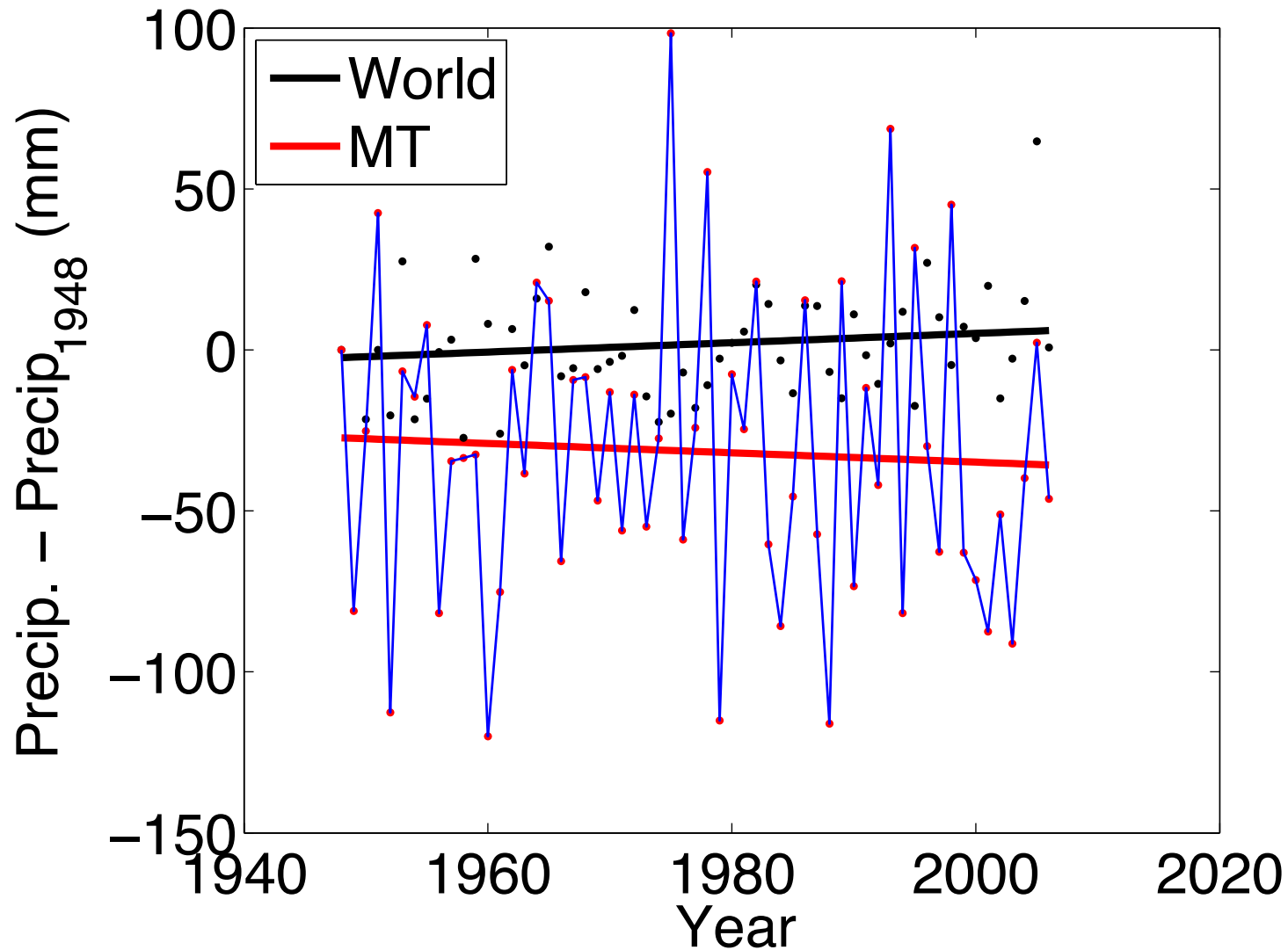
Defining areas belonging to MT from Sheffield et al.



MT *versus* global temperature trends

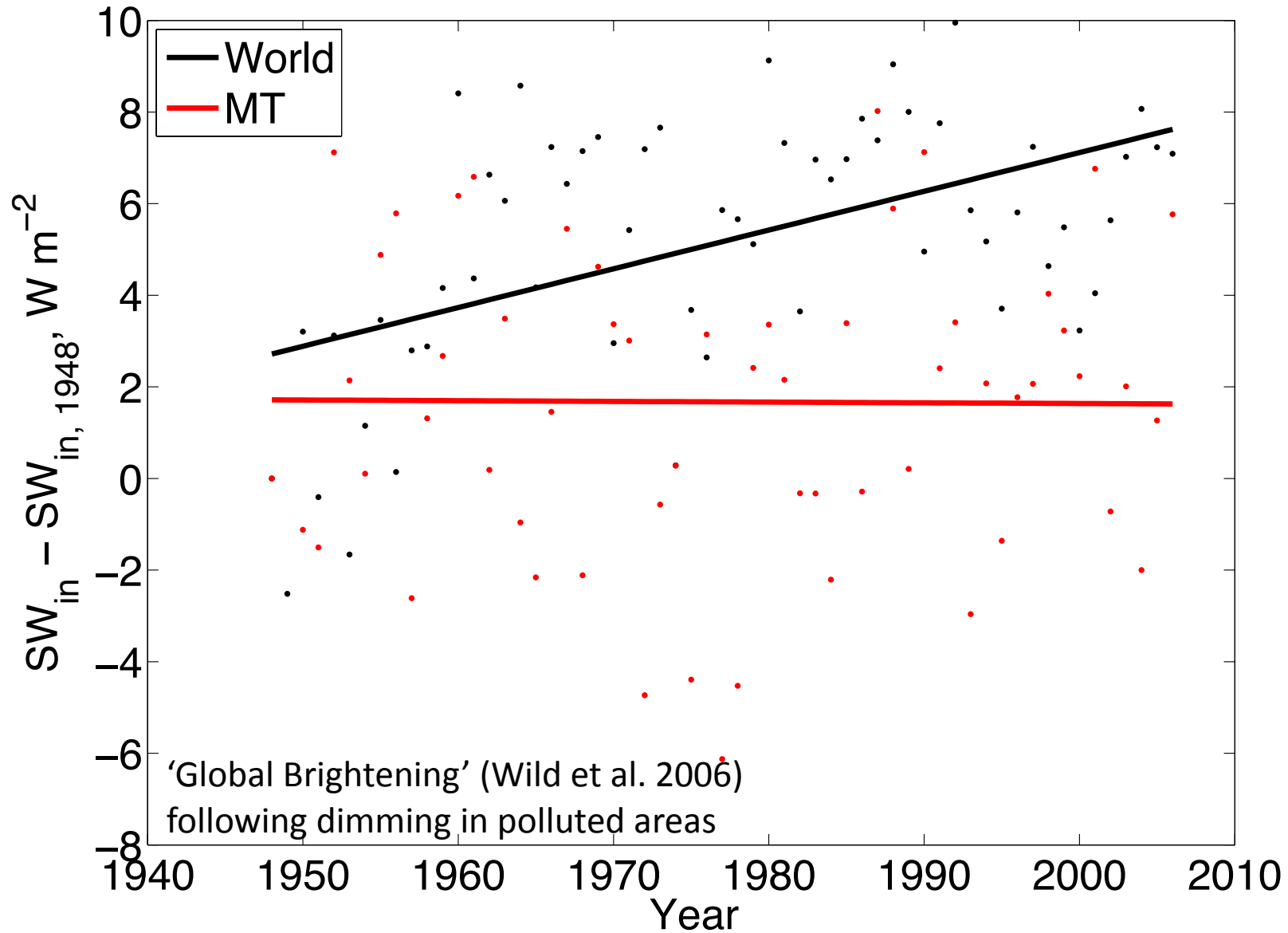


MT *versus* global precipitation trends

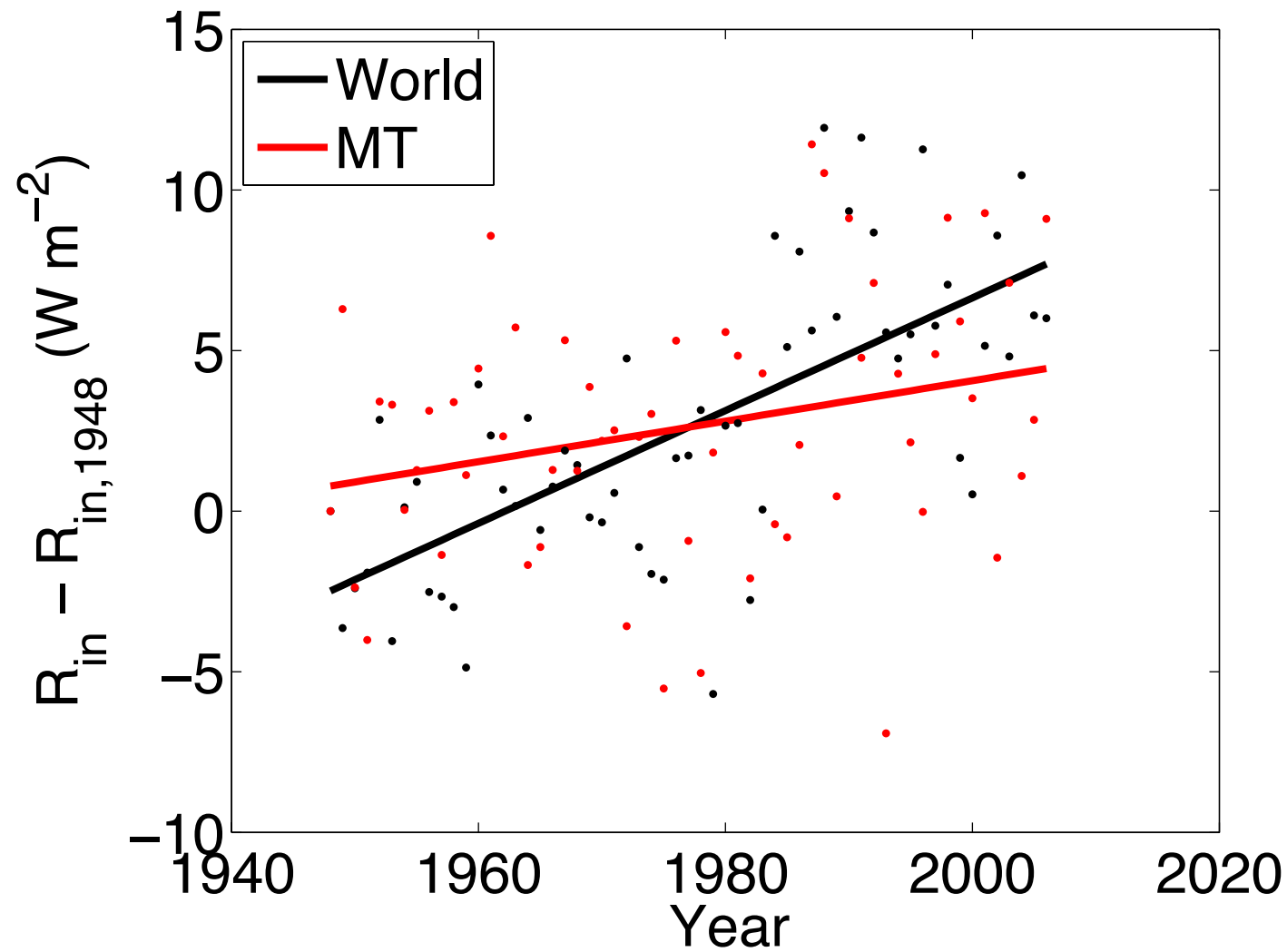


Sheffield et al. (2012) Little change in global drought over the past 60 years.
Nature doi:10.1038/nature11575

MT versus global incident shortwave trends



MT *versus* global incident radiation trends



Increases in aridity



Not Changing

Aridity index = Precipitation / Potential Evapotranspiration

Increasing as a function of incident radiation

Conclusion:

When making water use decisions we must recognize that there is now more energy to move water to the atmosphere.

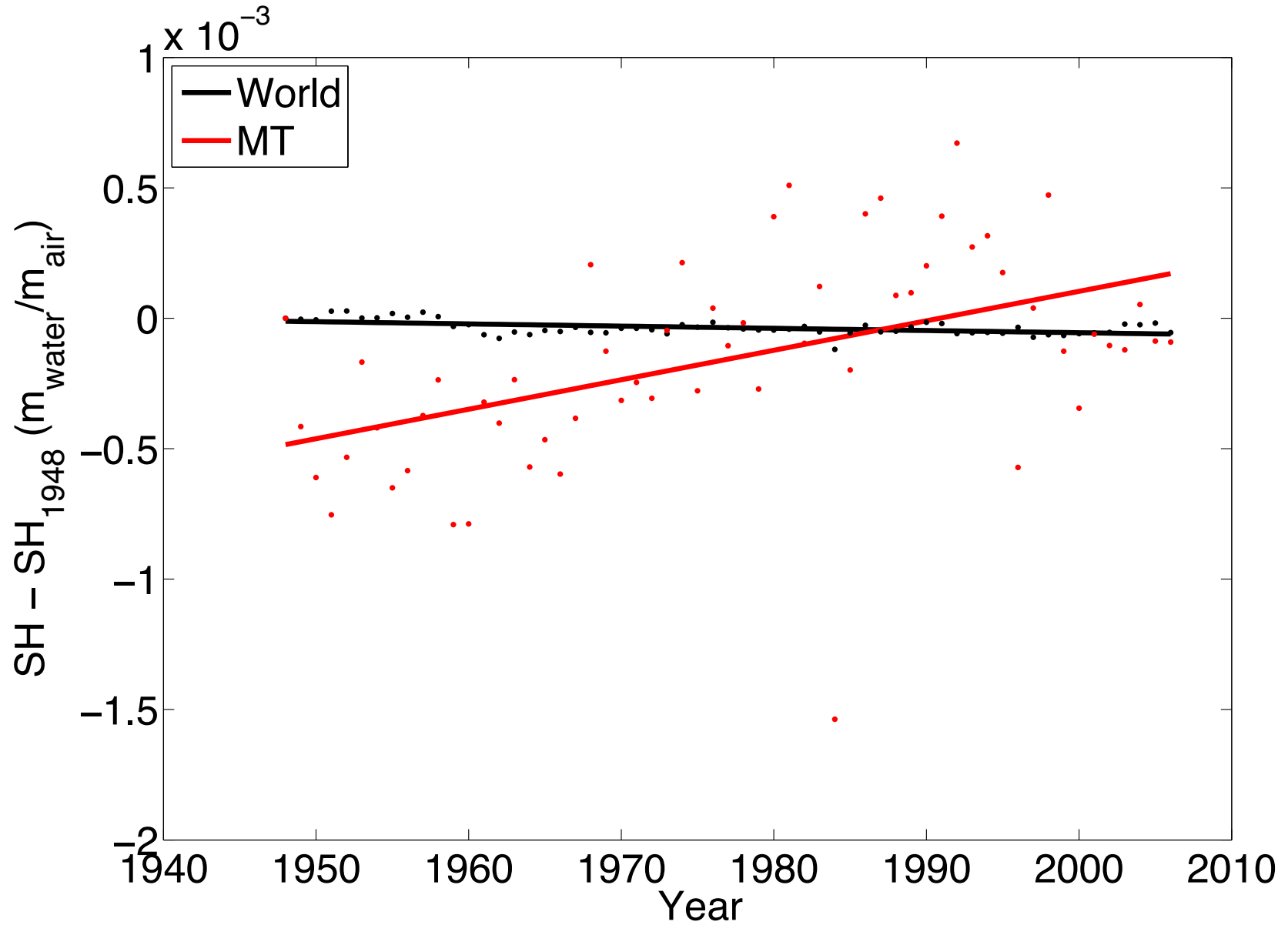
Trends in MT are not as acute as the global mean, but MT has less available water than the global mean.

Sheffield et al. data inputs



Dataset	Variables	Temporal coverage	Spatial coverage	Source
NCEP-NCAR reanalysis	P, T, SW, LW, q, Ps, w	1948-present, 6 hourly	Global, $\sim 2.0^\circ \times 2.0^\circ$	Kalnay et al. (1996)
CRU TS2.0	P, T, Cld	1901-2000, monthly	Global land excluding Antarctica, $0.5^\circ \times 0.5^\circ$	MCJHN
GPCP	P	1997-present, daily	Global, $1.0^\circ \times 1.0^\circ$	Huffman et al. (2001)
TRMM	P	Feb 2002-present, 3 hourly	$50^\circ S-50^\circ N, 0.25^\circ \times 0.25^\circ$	Huffman et al. (2003)
NASA Langley SRB	LW, SW	1983-95, monthly	Global, $1.0^\circ \text{ lat} \times 1.0^\circ-120^\circ \text{ lon}$	Stackhouse et al. (2004)

MT *versus* global specific humidity trends



The zero-dimensional climate model
*(Global change is simple, regional
change is difficult)*

=

Energy in

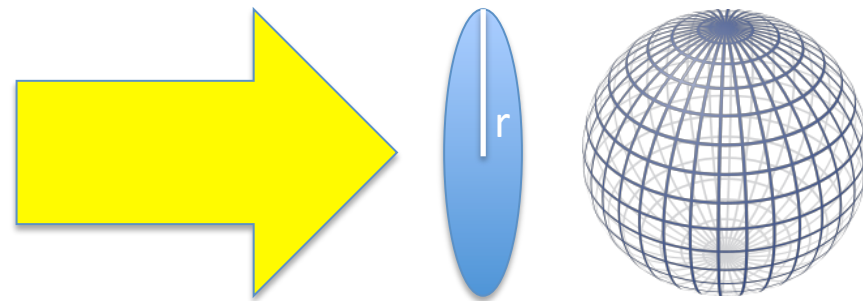


Energy out

At steady state

Energy in = solar energy

A disk of solar radiation of size πr^2 hits Earth



$$E_{in} = \pi r^2 S_c (1 - \rho)$$

$$A_{circle} = \pi r^2$$

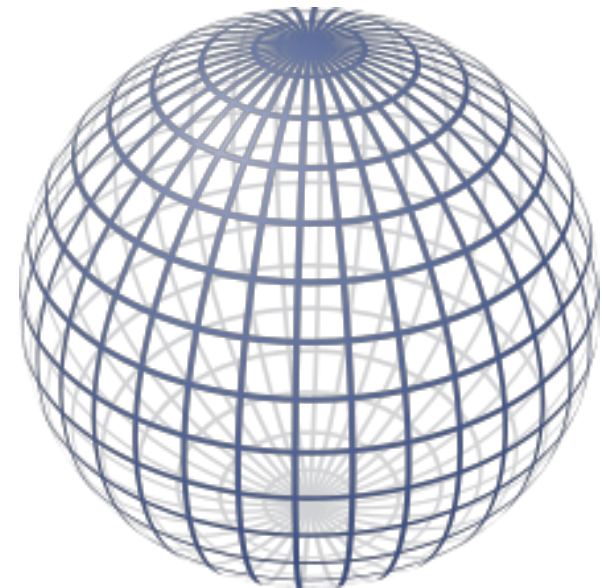
Energy out = planetary emittance

Stefan-Boltzmann Law

$$E_{out} = A\varepsilon\sigma T^4$$

$$E_{out} = 4\pi r^2\varepsilon\sigma T^4$$

$$A_{sphere} = 4\pi r^2$$



Planetary energy balance

Energy in = Energy out

$$E_{in} = E_{out}$$

$$\cancel{\pi r^2} S_c (1 - \rho) = 4 \cancel{\pi r^2} \epsilon \sigma T^4$$

$$\frac{S_c}{4} (1 - \rho) = \epsilon \sigma T^4$$

Solving for T

$$T^4 = \frac{S_c(1-\rho)}{4\varepsilon\sigma}$$

$$T = \sqrt[4]{\frac{S_c(1-\rho)}{4\varepsilon\sigma}}$$

Constants

$$\sigma = 5.67 \times 10^{-8} \text{ J K}^{-4} \text{ m}^{-2} \text{ s}^{-1}$$

Somewhat Constants

$$S_c = 1367 \text{ W m}^{-2}$$

$$\rho = 0.31$$

The temperature of blackbody Earth

$$T^4 = \frac{1367 \text{ J m}^{-2} \text{ s}^{-1} (1 - 0.31)}{4 \times 1 \times 5.67 \times 10^{-8} \text{ J K}^{-4} \text{ m}^{-2} \text{ s}^{-1}} = 254 \text{ K} = -19^\circ \text{ C}$$

The temperature of real Earth $\sim 14^\circ \text{ C}$

Why? The effective emissivity of Earth is less than 1.
This is the greenhouse effect

Solving for (effective) emissivity

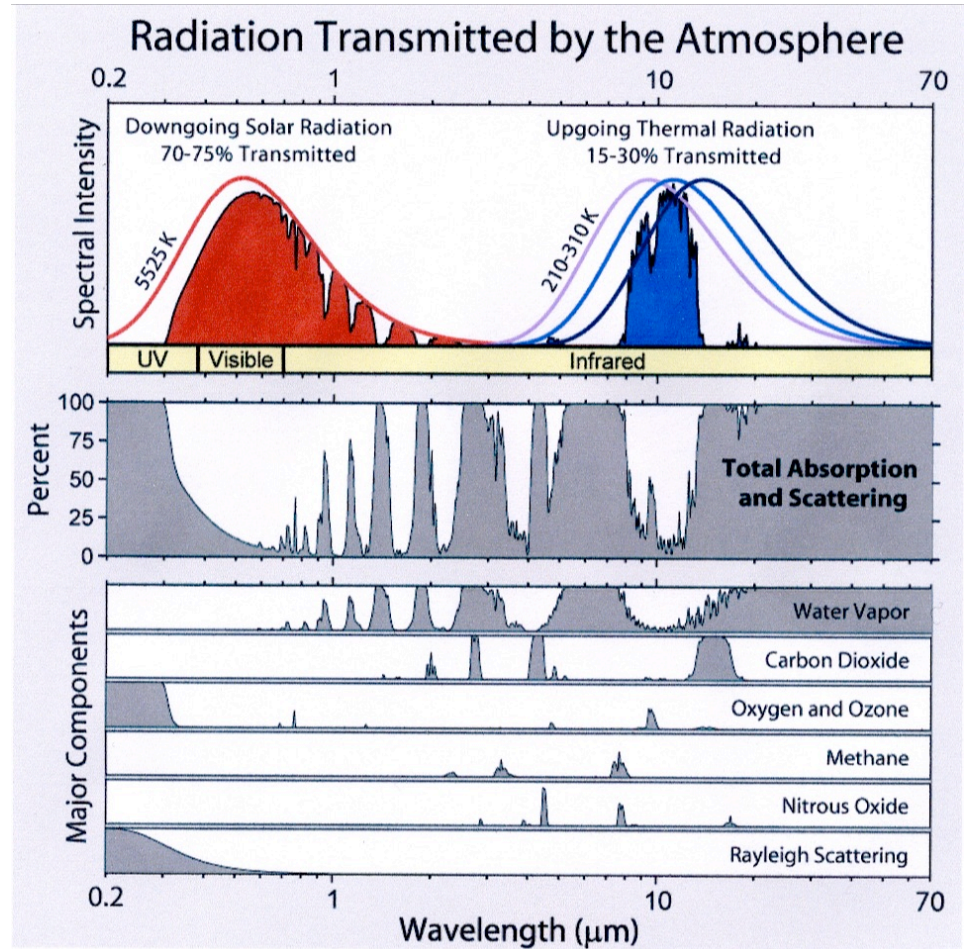
$$\varepsilon = \frac{S_c(1 - \rho)}{4\sigma T^4} \approx 0.6$$

Using $T = 14\text{ }^\circ\text{C}$

Kirchoff's Law of Radiation

$$\varepsilon(\lambda) \propto \alpha(\lambda)$$

Recall:



Albedo, emissivity and the solar constant determine global temperature

Why?
They determine
the energy balance.

